

# HDO and SO<sub>2</sub> thermal mapping on Venus

## VII. SO<sub>2</sub> variations with altitude and influence of the solar cycle

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Received 4 July 2025 / Accepted 17 October 2025

### ABSTRACT

*Context.* Sulfur dioxide and water are two key minor species of Venus' atmosphere which drive its chemical evolution. However, the long-term variations in the SO<sub>2</sub> abundance at the cloud top, measured since 1978, remain unexplained.

*Aims.* In order to address this question, since 2012, we have performed a ground-based campaign to monitor the SO<sub>2</sub> and H<sub>2</sub>O abundances in the region of the Venus upper cloud, using the TEXES (Texas Echelon Cross-Echelle Spectrograph) imaging spectrometer at the NASA InfraRed Telescope Facility (IRTF, Mauna Kea Observatory).

*Methods.* Observations were recorded in three spectral ranges: (1) the 1342–1348 cm<sup>-1</sup> (7.4 μm) spectral range, where SO<sub>2</sub>, CO<sub>2</sub> and HDO (used as a proxy for H<sub>2</sub>O) transitions are observed at an altitude of about 62 km, defined in our model as the cloud top; (2) the 529–530 cm<sup>-1</sup> range (18.9 μm), where SO<sub>2</sub> and CO<sub>2</sub> are probed within the clouds a few kilometers below the cloud top; (3) the 1160–1165 cm<sup>-1</sup> range (8.6 μm) where the weak SO<sub>2</sub> ν<sub>1</sub> band is used to probe a few kilometers above the cloud top.

*Results.* We present here the data recorded from July 2023 to February 2025. As was reported in our previous analyses, the SO<sub>2</sub> maps show evidence for the formation of SO<sub>2</sub> plumes, mostly located around the equator, with a typical lifetime of a few hours; large variations in the SO<sub>2</sub> disk-integrated abundance also appear on a timescale of a few months. In contrast, the H<sub>2</sub>O abundance is remarkably uniform over the disk and shows moderate variations as a function of time. The present dataset shows for the first time the detection of the SO<sub>2</sub> ν<sub>1</sub> band above the cloud top. The simultaneous analysis of the three SO<sub>2</sub> bands allows us to constrain the SO<sub>2</sub> volume mixing ratio within and above the cloud. In addition, we have re-analyzed the long-term evolution of the SO<sub>2</sub> and H<sub>2</sub>O mixing ratios at the cloud top. Between 2012 and 2023, the SO<sub>2</sub> abundance is clearly anti-correlated with the solar activity, which suggests that photochemistry, activated by the solar UV flux, is the main mechanism driving the SO<sub>2</sub> abundance. After 2023, SO<sub>2</sub> exhibits strong variations with a timescale as short as two months, which suggest that other mechanisms are involved. Finally, we have re-analyzed the distribution of SO<sub>2</sub> at the cloud top as a function of the local time, using three different data subsets (2012–2017, 2021–2022, 2023–2025). In spite of the differences, there is a general trend toward a minimum value around 12:00 and maxima around 5:00 and 17:00.

*Conclusions.* Our data suggest that the solar cycle plays an important role in the long-term evolution of SO<sub>2</sub> and H<sub>2</sub>O at the cloud top. In addition, in some cases, other mechanisms are probably also involved, possibly associated with dynamical motions, implying sublimation/condensation processes and/or gas-aerosol conversion.

**Key words.** planets and satellites: atmospheres – planets and satellites: composition – planets and satellites: terrestrial planets – planets and satellites: individual: Venus

## 1. Introduction

Water and sulfur dioxide are the drivers of Venus' atmospheric chemistry (Krasnopolsky 1986, Krasnopolsky 2007, 2010; Mills et al. 2007; Zhang et al. 2012). Below the clouds, both species are present with relatively large abundances (about 30 ppmv and 130 ppmv, respectively, Bézard and De Bergh 2012; Marcq et al. 2018) and, at low latitude, are transported upward by Hadley convection. Following the SO<sub>2</sub> photodissociation and the combination of SO<sub>3</sub> with H<sub>2</sub>O, sulfuric acid H<sub>2</sub>SO<sub>4</sub> is formed and condenses in the form of binary (H<sub>2</sub>SO<sub>4</sub>–H<sub>2</sub>O) aerosols in the cloud deck (Stolzenbach et al. 2023). Above the cloud top, the volume mixing ratios of H<sub>2</sub>O and SO<sub>2</sub> drop drastically to about

1–3 ppmv (Fedorova et al. 2008; Belyaev et al. 2012) and 10–1000 ppbv (Zasova et al. 1993; Marcq et al. 2013, 2020; Vandaele et al. 2017a,b), respectively. Since the condensation mechanism consumes 2 H<sub>2</sub>O molecules for one SO<sub>2</sub> molecule, another unknown mechanism is required to trap the extra SO<sub>2</sub> within the clouds.

Both species have been extensively monitored over several decades, with Pioneer Venus, the Venera 15 spacecraft, Venus Express and Akatsuki using imaging and spectroscopy in the ultraviolet and infrared ranges. As a complement to these datasets, since 2012, we have been using the TEXES (Texas Echelon Cross-Echelle Spectrograph) imaging spectrometer at the NASA InfraRed Telescope Facility (IRTF, Mauna Kea Observatory) to map SO<sub>2</sub> and H<sub>2</sub>O at the cloud top of Venus (using HDO as a proxy of H<sub>2</sub>O) and to monitor the behavior of these two

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species as a function of time. Eighteen runs have been recorded between January 2012 and July 2022 (Encrenaz et al. 2023, hereafter referred as E23, and references therein). The main result of these studies is that SO<sub>2</sub> and H<sub>2</sub>O exhibit very different behaviors. H<sub>2</sub>O is always uniformly distributed over the disk and shows moderate variations on the long term; in contrast, the SO<sub>2</sub> maps most often show sporadic plumes which appear and disappear with a typical timescale of a few hours; these features are probably the signature of dynamical activity within the clouds. The disk-integrated SO<sub>2</sub> abundance also shows strong variations over the long term, with a contrast factor of about 10 between the minimum value observed in February 2014 and the maximum value of July 2018 (also reobserved early 2025). A plausible explanation for the absence of H<sub>2</sub>O plumes is the fact that the H<sub>2</sub>O mixing ratio is globally constant within and above the clouds, so it is not a good tracer of vertical motions.

In this paper, we present the results obtained during the seven new observing runs recorded between March 2023 and March 2025. The observations are described in Section 2, showing in particular the detection of the SO<sub>2</sub>  $\nu_1$  band at 8.6  $\mu\text{m}$ . These data are used to constrain the SO<sub>2</sub> vertical profile at the three altitudes probed by the three SO<sub>2</sub> bands (Section 3). Then, we use the whole TEXES data set (2012–2022) at 7.4  $\mu\text{m}$  to study the long-term evolution of H<sub>2</sub>O and SO<sub>2</sub> (Section 4). Finally, we update our statistical analysis of the SO<sub>2</sub> plumes, regarding their appearance as a function of the local time (Section 5). Results and conclusions are presented in Section 6.

## 2. The data

TEXES (Texas Echelon Cross Echelle Spectrograph) is an imaging high-resolution thermal infrared spectrometer in operation at the NASA InfraRed telescope Facility (Lacy et al. 2002), which combines high spectral ( $R = 80\,000$  at 7  $\mu\text{m}$ ) and spatial (around 1 arcsec) capabilities. As for our previous observations, we selected three spectral ranges: (1) 1342–1348  $\text{cm}^{-1}$  (7.4  $\mu\text{m}$ ), (2) 529–530  $\text{cm}^{-1}$  (18.9  $\mu\text{m}$ ) and (3) 1160–1165  $\text{cm}^{-1}$  (8.6  $\mu\text{m}$ ). The first setting (7.4  $\mu\text{m}$ ) has been used since 2012 to monitor the mixing ratios of SO<sub>2</sub> and HDO at the altitude level  $z = 62$  km, referred in our model as the cloud top. The second setting (18.9  $\mu\text{m}$ ) has been also regularly used to monitor the SO<sub>2</sub> abundance a few kilometers below the cloud top ( $z = 57$  km). The third setting (8.6  $\mu\text{m}$ ) probes a few kilometers above the cloud top ( $z = 67$  km) and also contains the weak  $\nu_1$  band of SO<sub>2</sub>. This band is difficult to detect on Venus, both because of its intrinsic weakness and because the sharp depletion of SO<sub>2</sub> above the cloud top (E23). However, when the SO<sub>2</sub> abundance at the cloud top is over a few hundred ppbv, as in 2018 and 2019, the band might be detectable or its absence might bring a constraint on the SO<sub>2</sub> vertical gradient above the clouds. For this reason, we have observed this spectral range regularly since June 2021.

Seven runs have taken place in March, July, October, and December 2023, in February 2024, and in January and March 2025; the observing parameters are listed in Table 1. The length and the width of the slit were respectively 6.0 and 1.0 arcsec at 7.4  $\mu\text{m}$ , 12.0 and 2.0 arcsec at 18.9  $\mu\text{m}$ , and 8.0 and 1.0 arcsec at 8.6  $\mu\text{m}$ . As we did previously, we aligned the slit along the North-South celestial axis and we shifted it from west to east, with a step of half the slit width and an integration time of 2 seconds per position, to cover the planet in longitude from limb to limb, and to add a few pixels on the sky beyond each limb for sky subtraction. Since the rotation axis of Venus is close to the celestial axis, each scan corresponds to a given latitude range of

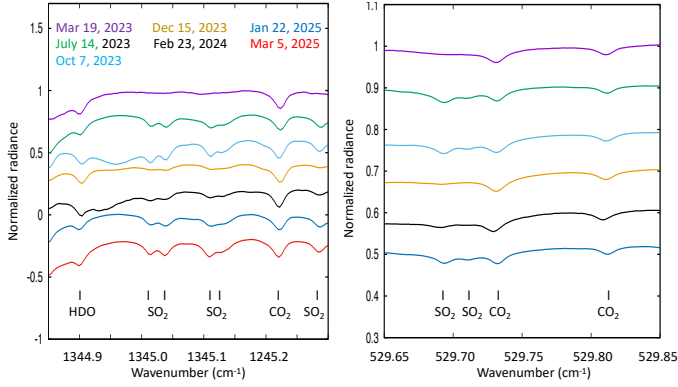
**Table 1.** Summary of TEXES runs in 2023–2025.

Date of obs.	Venus diameter (arcsec)	SEP longitude (degrees)	SEP 1 local time (hours)
2023/03/16–2023/03/21	13.0–13.2	360–13	15.3–15.5
2023/07/13–2023/07/19	40.0–44.5	294–304	20.2–20.8
2023/10/03–2023/10/08	31.0–29.0	53–65	5.0–5.3
2023/12/13–2023/12/18	15.8–15.3	233–247	7.7–7.9
2024/02/23–2024/02/26	11.3–11.1	65–82	9.5–9.7
2025/01/22–2025/01/25	28.0–29.1	244–251	18.4–18.6
2025/02/27–2025/03/05	47.3–51.6	319–328	20.9–21.5

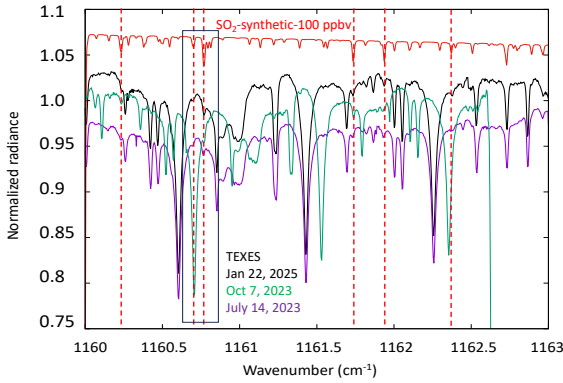
about 6 arcsec. As the diameter of Venus was always larger than the slit length, we multiplied the scans to cover the full latitude range from North to South with some overlap. The TEXES data cubes were calibrated using the standard radiometric method (Lacy et al. 2002; Rohlfis and Wilson 2004). Calibration frames consisting of three measurements (black chopper blade, sky and low-emissivity chopper blade) are systematically taken before each observing scan, and the difference (black-sky) is taken as a flat field. As the atmospheric correction, however, is not complete for all terrestrial atmospheric lines (partly because these lines are not all formed at the same atmospheric levels, and thus have different temperatures), the HDO, SO<sub>2</sub> and CO<sub>2</sub> lines selected for our study are located outside the atmospheric features.

Figure 1 shows representative disk-integrated spectra corresponding to the 2023–2025 runs in the 7.4 and 18.9  $\mu\text{m}$  ranges. The 7.4- $\mu\text{m}$  spectral range (1344.8–1345.4  $\text{cm}^{-1}$ ) includes several weak SO<sub>2</sub> transitions, two weak CO<sub>2</sub> lines and one weak HDO line. As in our previous studies (see E23 and references therein), we used the HDO line at 1344.90  $\text{cm}^{-1}$ , the SO<sub>2</sub> multiplet at 1345.1  $\text{cm}^{-1}$  and the CO<sub>2</sub> line at 1345.22  $\text{cm}^{-1}$ . In the 18.9  $\mu\text{m}$  range (529–530  $\text{cm}^{-1}$ ) weak transitions of SO<sub>2</sub> and CO<sub>2</sub> are found. We used the SO<sub>2</sub> multiplet at 529.69  $\text{cm}^{-1}$  and the CO<sub>2</sub> line at 529.81  $\text{cm}^{-1}$ . It immediately appears, in both datasets, that the SO<sub>2</sub> lines depths show strong variations from run to run: they are as strong as the neighbor CO<sub>2</sub> line at maximum, and almost disappear at minimum. In order to estimate the H<sub>2</sub>O and SO<sub>2</sub> mixing ratios, as was discussed in earlier papers (see e.g., E23), we use the ratios of the SO<sub>2</sub>/CO<sub>2</sub> and HDO/CO<sub>2</sub> line depths which provide a good approximation of the SO<sub>2</sub> and H<sub>2</sub>O mixing ratios at the altitude probed in each spectral range. The HDO/H<sub>2</sub>O ratio at the cloud deck and above is estimated to 200 times the terrestrial value VSMOW (Fedorova et al. 2008; E23).

In addition to the 7.4 and 18.9  $\mu\text{m}$  spectral ranges, since 2021, we have observed the 1160–1165  $\text{cm}^{-1}$  range, at 8.6  $\mu\text{m}$ , which contains some transitions of the SO<sub>2</sub>  $\nu_1$  band at 8.6  $\mu\text{m}$  (Zasova et al. 1993). As mentioned above, its detection is difficult. We failed to detect it in 2021 and 2022, which was not surprising as the SO<sub>2</sub> abundance at that time was especially low. As shown in Figure 1, the SO<sub>2</sub> abundance has been very high in July 2023, October 2023 and January 2025. As a consequence, the  $\nu_1$  band



**Fig. 1.** Examples of disk-integrated spectra of Venus (normalized radiance) recorded in 2023–2025. Left: 1344.85–1345.30  $\text{cm}^{-1}$  spectral range (7.4  $\mu\text{m}$ ). The spectra are shifted by  $-0.2$  in vertical scale for clarity. Right: examples of disk-integrated spectra of Venus (normalized radiance) between 529.65 and 529.85  $\text{cm}^{-1}$  (18.9  $\mu\text{m}$ ). The spectra are shifted by  $-0.1$  in vertical scale for clarity.



**Fig. 2.** Disk-integrated spectra of Venus between 1160 and 1163  $\text{cm}^{-1}$  (8.6  $\mu\text{m}$ ) recorded on July 14, 2023, October 7, 2023 and January 22, 2025. The positions of the strongest  $\text{SO}_2$  lines are indicated in red by vertical dashed lines. The terrestrial lines coincide in the July 2023 and January 2025 because the Doppler shift is the same for both datasets, while it is different for the spectrum taken in October 2023. The spectra are dominated by terrestrial atmospheric absorption features, mostly  $\text{N}_2\text{O}$  and  $\text{O}_3$ . The black rectangle shows the spectral range used for our analysis.

has been actually detected in each of these three observing runs. This is its first ground-based detection; the only earlier report was in 1983 by the Venera 15 far infrared interferometer, in some specific locations on the planet (Zasova et al. 1993).

Figure 2 shows three disk-integrated spectra of Venus recorded in the 1160–1163  $\text{cm}^{-1}$  spectral range, in July 2023, October 2023 and January 2025, respectively. The spectrum is fully dominated by terrestrial atmospheric lines, mostly due to nitrous oxide and ozone. As expected, the  $\text{SO}_2$  lines are very weak, and were not detected in March 2023, December 2023 and February 2024. We choose to study the spectral region around the  $\text{SO}_2$  transition at 1160.7733  $\text{cm}^{-1}$  ( $I = 5.57 \times 10^{-21}$   $\text{cm}^2/\text{mol}$ ,  $E = 74.66$   $\text{cm}^{-1}$ ) which is well separated from terrestrial lines whatever the Doppler shift is (Fig. 2).

### 3. The $\text{SO}_2$ vertical distribution

As was discussed in our earlier publications (see e.g., E23), observing Venus at three different infrared wavelengths allows us

to probe different atmospheric levels, because of the variations in the  $\text{H}_2\text{SO}_4$  extinction coefficient that is maximum around 8.6  $\mu\text{m}$  and minimum around 19  $\mu\text{m}$ , while its value at 7.4  $\mu\text{m}$  is intermediate. Based on a thermal profile inferred from  $\text{CO}_2$  TEXES observations (Giles et al. 2022), our previous analyses of the  $\text{CO}_2$  lines at 7.4 and 18.9  $\mu\text{m}$  have shown that the penetration levels at these two wavelengths are located near  $P = 250$  mbar ( $T = 241$  K,  $z = 57$  km) at 18.9  $\mu\text{m}$  and  $P = 100$  mbar ( $T = 230$  K,  $z = 62$  km, referred as the cloud top in our model) at 7.4  $\mu\text{m}$ . Note that this definition of the cloud top is somehow arbitrary: it is the penetration level ( $\tau = 1$ ) at 7.4  $\mu\text{m}$  and is slightly model dependent, as the  $\text{CO}_2$  line profiles give access to the pressure, and not the altitude; in addition, this level shows slight variations from run to run. We note that our reference level is a few km lower than the cloud top defined at this wavelength by Titov et al. (2018) for the standard model, and better fits the model with enriched upper haze.

In our earlier work, we have inferred the  $\text{SO}_2$  mixing ratio at  $z = 57$  and 62 km from the analysis of the  $\text{SO}_2$  lines. The detection of the  $\text{SO}_2$  lines at 8.6  $\mu\text{m}$  gives us a new constrain on the  $\text{SO}_2$  vertical distribution above the cloud top. From the shape of the  $\text{SO}_2$  lines (broader than the  $\text{CO}_2$  lines) at 7.4 and 18.9  $\mu\text{m}$ , it could be seen that depletion of  $\text{SO}_2$  had to occur a few kilometers above the cloud top. In the model described in E23, we introduced this cutoff at  $z = 70$  km ( $T = 204$  K,  $P = 23$  mbar). Such a strong depletion with altitude is also observed in the UV range (Marcq et al. 2020). Using our radiative transfer calculations, we derived the  $\text{H}_2\text{O}$  and  $\text{SO}_2$  mixing ratios from the line depth ratios of weak  $\text{H}_2\text{O}$  and  $\text{SO}_2$  lines divided by a weak nearby  $\text{CO}_2$  line, as we did in our earlier studies (E23 and references therein). This simple method has the advantage of providing a good estimate of the  $\text{H}_2\text{O}$  and  $\text{SO}_2$  mixing ratios, independently of the atmospheric parameters and the airmass factor. We have been using it to study the long-term evolution of the two species at the cloud top, and to obtain maps of the  $\text{H}_2\text{O}$  and  $\text{SO}_2$  mixing ratios at 7.3 and 18.9  $\mu\text{m}$ . For the conversion from the line depth ratios ( $\text{ldr}$ ) into the volume mixing ratios ( $\text{vmr}$ ), the following equations are used (E23):

At 7.4  $\mu\text{m}$  (1345  $\text{cm}^{-1}$ ):

$$\text{vmr}(\text{SO}_2)(\text{ppbv}) = \text{ldr}(\text{SO}_2) \times 600.0$$

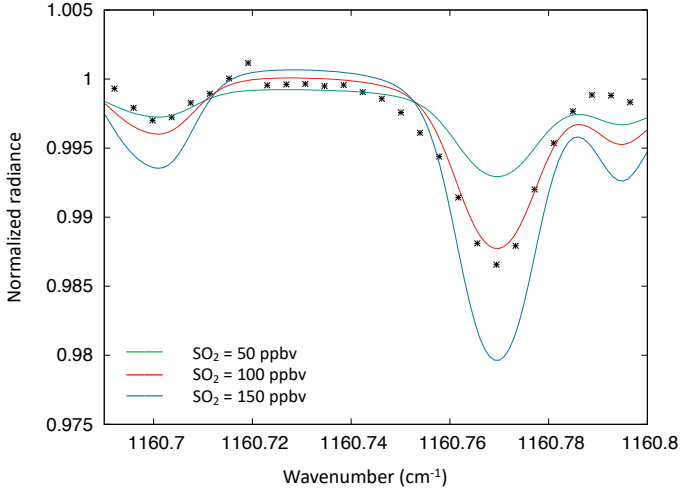
$$\text{vmr}(\text{H}_2\text{O})(\text{ppmv}) = \text{ldr}(\text{H}_2\text{O}) \times 1.5.$$

At 19  $\mu\text{m}$  (530  $\text{cm}^{-1}$ ):

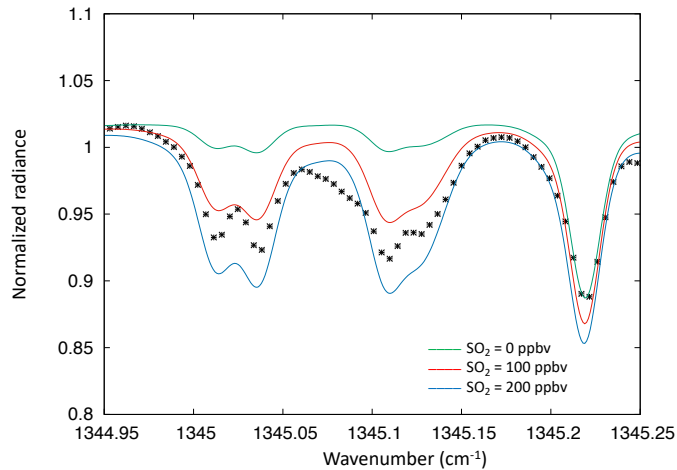
$$\text{vmr}(\text{SO}_2)(\text{ppbv}) = \text{ldr}(\text{SO}_2) \times 500.0.$$

In the case of the 8.6- $\mu\text{m}$  spectra, we cannot apply this method, unfortunately, because there is no  $\text{CO}_2$  transition available in this spectral range. The only information about the penetration level at 8.6  $\mu\text{m}$  comes from the infrared spectrum of the  $\text{H}_2\text{SO}_4$  absorption coefficient (Zasova et al. 1993) which indicates that it is higher at 8.6  $\mu\text{m}$  than at 7.4  $\mu\text{m}$ . Then, we deduce that the penetration level at 8.6  $\mu\text{m}$  must be a few km higher than the one at 7.4  $\mu\text{m}$ . We modeled the  $\text{SO}_2$  line around 1160.77  $\text{cm}^{-1}$  assuming different altitude levels and different values of the  $\text{SO}_2$   $\text{vmr}$  above the cloud top and trying to fit the three disk-integrated spectra simultaneously. Our radiative transfer code is described in earlier publications (E23 and references therein). It is a forward model and no retrieval algorithm is used in our study. Our model consists in a line-by-line calculation without scattering. Spectroscopic data are taken from the GEISA data base (Jacquinet-Husson et al. 2018). In the case of  $\text{SO}_2$ , for which the broadening coefficients by  $\text{CO}_2$  are not given in the database, as in our earlier studies, we multiplied the broadening coefficients by  $\text{N}_2$  by a factor 1.4 (Nakazawa and Tanaka 1982).

On the basis of the equations mentioned above, the  $\text{SO}_2$   $\text{vmr}$  was set to 500 ppbv at 57 km and 400 ppbv at 62 km for the three

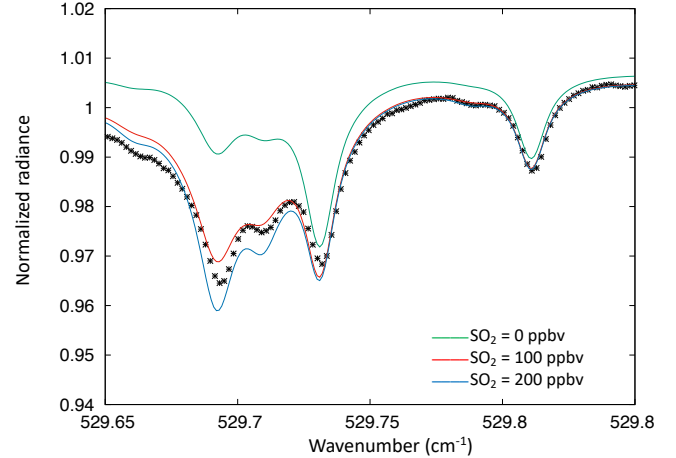


**Fig. 3.** Disk-integrated spectrum of Venus between 1160.69 and 1160.80  $\text{cm}^{-1}$  (8.6  $\mu\text{m}$ ) recorded on July 14, 2023. Models:  $\text{SO}_2(57 \text{ km}) = 500$  ppbv,  $\text{SO}_2(62 \text{ km}) = 400$  ppbv,  $\text{SO}_2(67 \text{ km}) = 50$  ppbv (green), 100 ppbv (red) and 150 ppbv (blue).

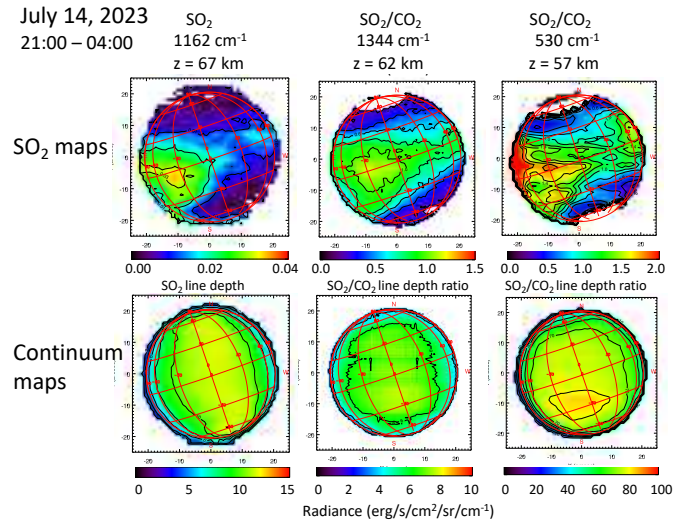


**Fig. 4.** Disk-integrated spectrum of Venus between 1344.95 and 1345.25  $\text{cm}^{-1}$  (7.4  $\mu\text{m}$ ) recorded on July 14, 2023. Models:  $\text{SO}_2(57 \text{ km}) = 500$  ppbv,  $\text{SO}_2(62 \text{ km}) = 400$  ppbv,  $\text{SO}_2(67 \text{ km}) = 0$  ppbv (green), 100 ppbv (red) and 200 ppbv (blue).

data sets showing the  $\text{SO}_2$  line at 8.6  $\mu\text{m}$ . An example of fit is shown in Figures 3–5. In this model, a cutoff is applied at  $z = 72$  km ( $T = 201$  K,  $P = 13$  mbar). The influence of this parameter on the results is discussed in Section 5. From the direct comparison between the data and the models, in Figures 3–5, it can be seen that, at 7.4 and 18.9  $\mu\text{m}$ , the best fit is obtained with a  $\text{SO}_2$  vmr of 150 ppbv, while the 8.6  $\mu\text{m}$  spectrum favors a  $\text{SO}_2$  vmr of 100 ppbv. This discrepancy probably means that the choice of the  $\text{SO}_2$  cutoff altitude (fixed, in our earlier studies, from the shapes of the  $\text{SO}_2$  lines at 7.4 and 18.9  $\mu\text{m}$ ) is not optimized. It also probably illustrates that our model is too simple and does not account, in particular, for possible inhomogeneities over the Venus disk. From Figures 3–5, we conclude that a  $\text{SO}_2$  vmr between 100 and 150 ppbv at an altitude of 67 km ( $T = 223$  K,  $P = 40$  mbar), associated with a  $\text{SO}_2$  cutoff at  $z = 72$  km, provides a reasonable fit of the three spectra. It must be reminded that this solution is not unique, as the  $\text{SO}_2$  measurement at 8.7  $\mu\text{m}$  only provides a constraint upon the  $\text{SO}_2$  column density. The  $\text{SO}_2$  vmr estimated at



**Fig. 5.** Disk-integrated spectrum of Venus between 529.65 and 529.85  $\text{cm}^{-1}$  (18.9  $\mu\text{m}$ ) recorded on July 14, 2023. Models:  $\text{SO}_2(57 \text{ km}) = 500$  ppbv,  $\text{SO}_2(62 \text{ km}) = 400$  ppbv,  $\text{SO}_2(67 \text{ km}) = 0$  ppbv (green), 100 ppbv (red) and 200 ppbv (blue).

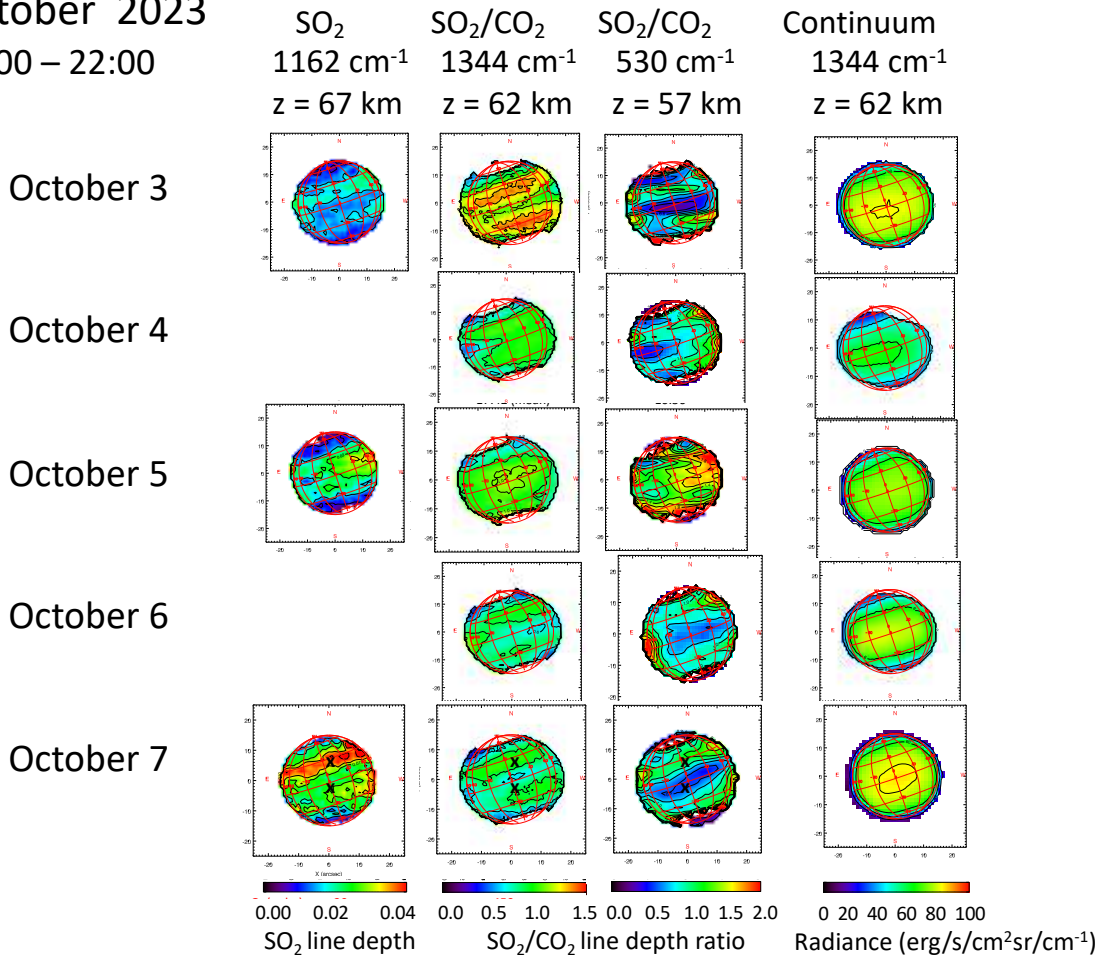


**Fig. 6.** Top:  $\text{SO}_2$  maps recorded on July 14, 2023. Left:  $z = 67$  km ( $\text{SO}_2$  line depth); middle:  $z = 62$  km ( $\text{SO}_2/\text{CO}_2$  line depth ratio); right:  $z = 57$  km ( $\text{SO}_2/\text{CO}_2$  line depth ratio). Disk-integrated spectra corresponding to these maps are shown in Figs. 3–5. Bottom: continuum maps recorded on July 14, 2023 at  $z = 67$  km (left), 62 km (middle) and 57 km (right).

8.7  $\mu\text{m}$  thus depends on both the altitude of the penetration level (here 67 km) and the altitude of the  $\text{SO}_2$  cutoff (here  $z = 72$  km).

As the 8.6  $\mu\text{m}$  spectral range does not include  $\text{CO}_2$  transitions, we cannot obtain  $\text{SO}_2$  maps at 67 km, as we did at 57 and 62 km using the 18.9  $\mu\text{m}$  and 7.4  $\mu\text{m}$  data. We cannot either determine precisely the penetration level at 8.6  $\mu\text{m}$ . Still, we can use the depth of the  $\text{SO}_2$  transition at 1160.77  $\text{cm}^{-1}$  as an indicator of the  $\text{SO}_2$  abundance a few kilometers above our reference level probed at 7.4  $\mu\text{m}$  (in our model,  $z = 62$  km). Figures 6–8 show these maps, compared to the maps of the  $\text{SO}_2$  volume mixing ratio obtained at 7.4 and 19  $\mu\text{m}$ , from the data shown in Figure 1, using the transitions mentioned above. It can be seen that, not surprisingly, the  $\text{SO}_2$  spatial distribution at 62 km is usually well correlated with those observed at deeper levels. However, a remarkable behavior is observed in October 2023 (Fig. 7). As was discussed in earlier papers, the distribution of the  $\text{SO}_2$  plumes is usually concentrated around the equator

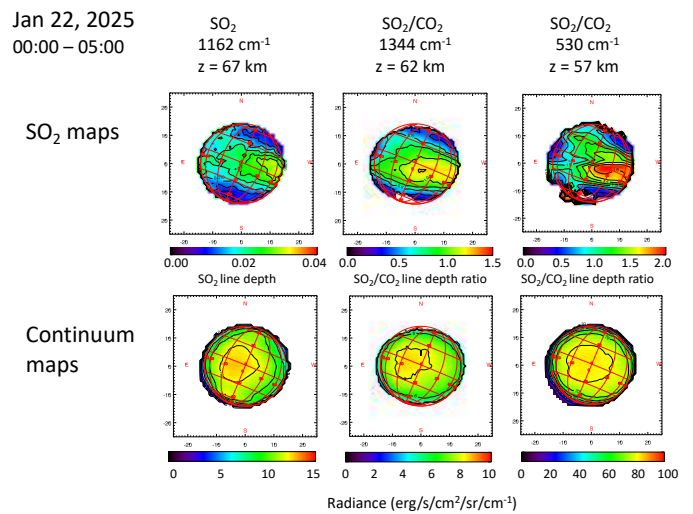
October 2023  
16:00 – 22:00



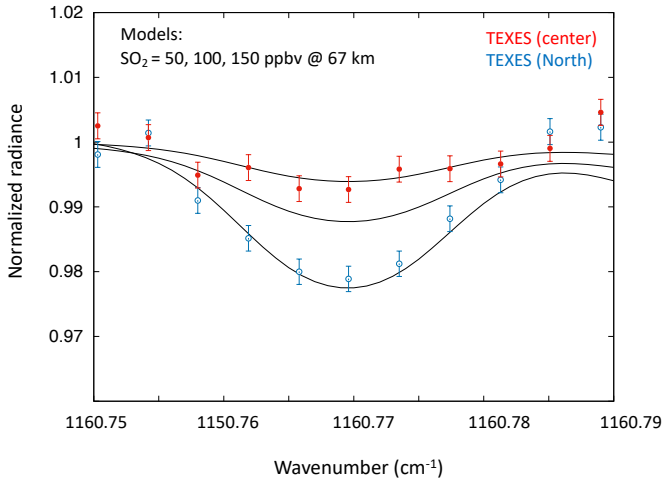
**Fig. 7.** SO<sub>2</sub> maps recorded on October 3–7, 2023. Left: z = 67 km (SO<sub>2</sub> line depth); middle: z = 62 km (SO<sub>2</sub>/CO<sub>2</sub> line depth ratio); next: z = 57 km (SO<sub>2</sub>/CO<sub>2</sub> line depth ratio). Last right column: continuum maps recorded on October 3–7, 2023 at z = 62 km. The black crosses on the October 7 SO<sub>2</sub> maps indicate the positions corresponding to the spectra shown in Figs. 9–11 (disk center and 30°N latitude).

(E23). In October 2023, the SO<sub>2</sub> plumes are structured in two parallel bands located at mid-latitude, over a longitude range of about 120 degrees. This band structure appears on October 3 and is observed over 4 consecutive days; indeed, as shown by the different 18.9 μm SO<sub>2</sub> maps between October 3 and October 7, it rotates by 90 degrees per day, following the 4-day rotation of the clouds. The structure also appears, although less clearly, on the SO<sub>2</sub> maps at higher altitudes; on October 7, the SO<sub>2</sub> map at z = 67 km shows clearly the band structure with a strong enhancement in the northern hemisphere. The longevity of this dynamical structure is actually exceptional, as the life time of SO<sub>2</sub> plumes are usually shorter than a day (E23). It can be noticed also that the parallel latitudinal structure is not observed on the continuum maps which reflect the temperature field at the three altitudes probed in this study.

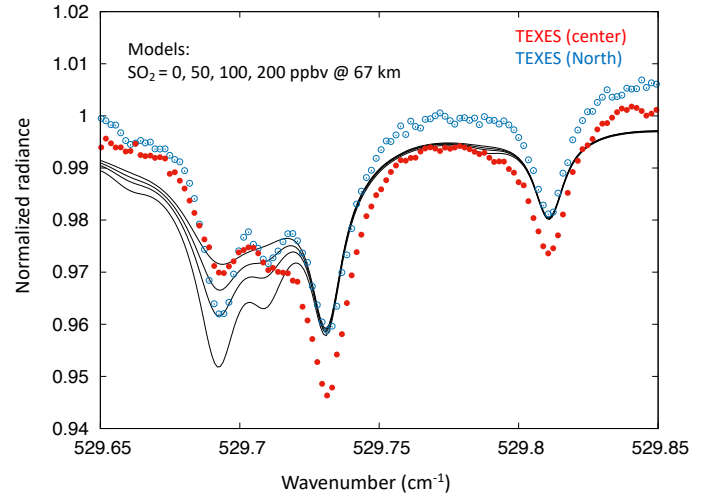
In order to investigate in more detail the anomaly observed in October 2023, we have selected, on the October 7 map, two locations close to the central meridian, at the disk center and at mid-northern latitude (Fig. 7). Figures 9–11 show the spectra of these two regions in the three spectral ranges. As expected from Figure 7, very strong variations are observed in the SO<sub>2</sub> absorptions between the center and the northern hemisphere. In particular, the very low SO<sub>2</sub> abundance observed at 18.9 μm



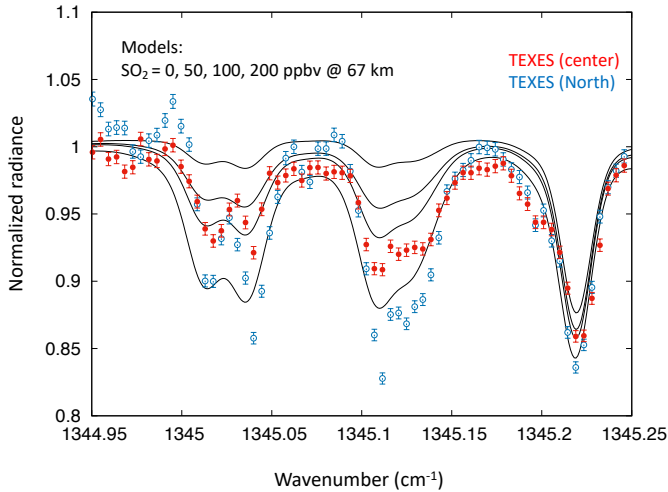
**Fig. 8.** Top: SO<sub>2</sub> maps recorded on January 22, 2025. Left: z = 67 km (SO<sub>2</sub> line depth); middle: z = 62 km (SO<sub>2</sub>/CO<sub>2</sub> line depth ratio); right: z = 57 km (SO<sub>2</sub>/CO<sub>2</sub> line depth ratio). Bottom: continuum maps recorded on January 22, 2025 at z = 67 km (left), 62 km (middle) and 57 km (right).



**Fig. 9.** Spectra of Venus between  $1160.75$  and  $1160.79$   $\text{cm}^{-1}$  ( $8.6$   $\mu\text{m}$ ) recorded on October 7, 2023 at the disk center (red points) and at  $30^\circ\text{N}$  latitude (blue points). Models:  $\text{SO}_2(57$  km) =  $500$  ppbv,  $\text{SO}_2(62$  km) =  $400$  ppbv. From top to bottom:  $\text{SO}_2(67$  km) =  $50$ ,  $100$  and  $200$  ppbv.



**Fig. 11.** Spectra of Venus between  $529.65$  and  $529.85$   $\text{cm}^{-1}$  ( $18.9$   $\mu\text{m}$ ) recorded on October 7, 2023 at the disk center (red points) and at  $30^\circ\text{N}$  latitude (blue points). Models:  $\text{SO}_2(57$  km) =  $500$  ppbv,  $\text{SO}_2(62$  km) =  $400$  ppbv. From top to bottom:  $\text{SO}_2(67$  km) =  $0$ ,  $50$ ,  $100$  and  $200$  ppbv.



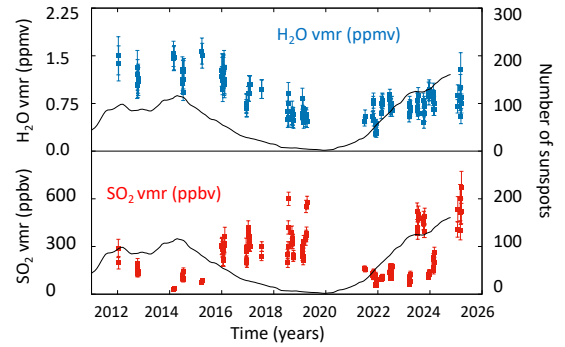
**Fig. 10.** Spectra of Venus between  $1344.95$  and  $1345.25$   $\text{cm}^{-1}$  ( $7.4$   $\mu\text{m}$ ) recorded on October 7, 2023 at the disk center (red points) and at  $30^\circ\text{N}$  latitude (blue points). Models:  $\text{SO}_2(57$  km) =  $500$  ppbv,  $\text{SO}_2(62$  km) =  $400$  ppbv. From top to bottom:  $\text{SO}_2(67$  km) =  $0$ ,  $50$ ,  $100$  and  $200$  ppbv.

( $z = 57$  km) at the disk center seems difficult to reconcile with the two other spectra; this discrepancy is discussed in Section 6.

#### 4. Long-term variations in $\text{SO}_2$ and $\text{H}_2\text{O}$

Figure 12 shows the temporal variations in the disk-integrated  $\text{H}_2\text{O}$  and  $\text{SO}_2$  mixing ratios at the cloud top ( $z = 62$  km) derived from the  $7.4$ - $\mu\text{m}$  data. The evolution of the solar activity is shown for comparison.

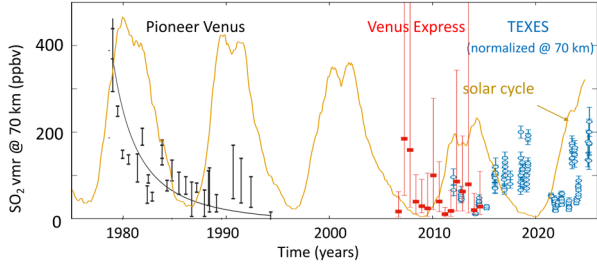
In our previous analysis (E23), we reported a clear anti-correlation of the two molecules (correlation coefficient =  $-0.9$ ) between 2014 and 2019. Since this period,  $\text{H}_2\text{O}$  and  $\text{SO}_2$  have evolved differently, with strong temporal variations in the  $\text{SO}_2$  abundance (by a factor as much as 5), over timescales as short as two months, while the  $\text{H}_2\text{O}$  abundance has remained more or less constant. It is interesting to note that, between 2012 and 2023, the  $\text{SO}_2$  abundance is anti-correlated with the solar activity, while the  $\text{H}_2\text{O}$  abundance shows a correlation with the solar



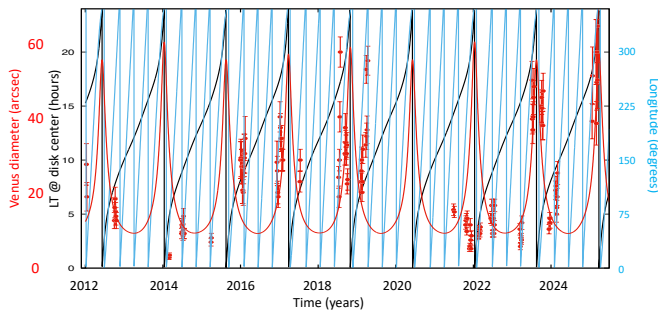
**Fig. 12.** Top: long-term variations in the disk-integrated  $\text{H}_2\text{O}$  volume mixing ratio (blue points), inferred from the HDO measurements at the cloud top from the TEXES data at  $7.4$   $\mu\text{m}$ . Bottom: long-term variations in the disk-integrated  $\text{SO}_2$  volume mixing ratio inferred at the cloud top ( $7.4$   $\mu\text{m}$ , red points). A daily value is shown in this figure. The error bars represent the uncertainty in the disk-integrated HDO and  $\text{SO}_2$  volume mixing ratios, averaged over a day. They depend upon the quality of the terrestrial atmospheric transparency, the size of the planetary disk, the continuum level and the number of maps recorded each day. In both cases, the solar activity (measured by the sunspot number) is shown for comparison.

cycle between 2012 and 2023. This strongly suggests that, in this period, photochemistry is the main mechanism driving the behaviors of  $\text{SO}_2$  and  $\text{H}_2\text{O}$  at the cloud top of Venus. Indeed, an increase in the solar UV flux is expected to enhance the photodissociation of  $\text{SO}_2$ , decreasing its abundance. As shown by Parkinson et al. (2015),  $\text{H}_2\text{O}$  and  $\text{SO}_2$  are expected to regulate themselves through the formation of  $\text{H}_2\text{SO}_4$  and this effect could be responsible for the anti-correlation observed between the two molecules.

We then wondered if the effect of the solar cycle could have been visible in the past. There is no long-term record of the water abundance at the cloud top (as no strong variations have been observed), but the  $\text{SO}_2$  abundance has been measured since 1978, thanks to the UV measurements of Pioneer Venus, and later Venus Express and Akatsuki. Figure 13 shows the long-term evolution of  $\text{SO}_2$  from 1978 to now, compared to the solar cycle. Because the UV data probe an altitude slightly higher than



**Fig. 13.** Long-term variations in the  $\text{SO}_2$  volume mixing ratio measured at the cloud top between 1978 and 2025. The data prior to 2012 are taken from [Esposito et al. 1984](#) (Pioneer Venus) and from [Marcq et al. 2020](#) (Venus Express). They have been obtained in the UV, at an altitude of about 70 km, and the TEXES data have been normalized accordingly. The solar activity (measured by the sunspot number) is shown for comparison.

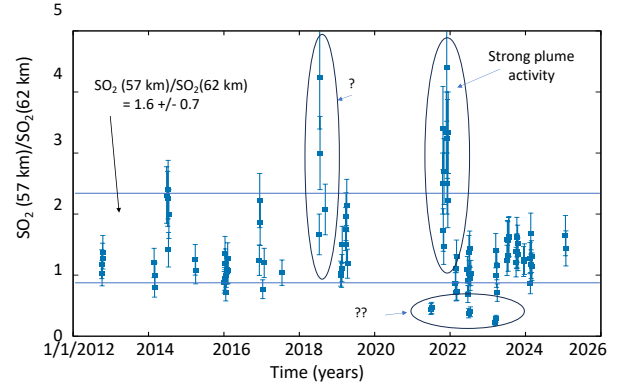


**Fig. 14.** Long-term variations in the  $\text{SO}_2$  volume mixing ratio measured by TEXES at the cloud top between 2012 and 2025 (red points). The Venus diameter (red line), the local time at the disk center (black line) and the longitude at the disk center (blue line) are shown for comparison.

the TEXES data (around  $z = 70$  km), we normalized our data by adjusting them to the Venus Express data taken in 2012–2015 by dividing them by a factor 3, which is consistent with the  $\text{SO}_2$  vmr inferred from our data at  $8.6 \mu\text{m}$ , as was discussed above. It can be seen that the strong decrease of the  $\text{SO}_2$  abundance observed by Pioneer Venus in 1978–1982 can be explained by the strong increase in the solar UV flux during these periods.

However, the anti-correlation between the  $\text{SO}_2$  abundance and the solar activity is not observed at all times. In particular, the strong  $\text{SO}_2$  decrease observed by Venus Express in 2006–2009 corresponds to a minimum of the solar activity; in addition, no increase in  $\text{SO}_2$  is observed in 1985–1987 when the solar activity is at its minimum. This suggests that the solar cycle is not the only driver of the  $\text{SO}_2$  long-term evolution. This conclusion is also supported by the TEXES data after 2023, suggesting that a change of regime has occurred at that time. The strong fluctuations of the  $\text{SO}_2$  abundances, on timescales of only two months, cannot be explained by the solar cycle. In order to study which parameter could be responsible for these fluctuations, we have plotted the  $\text{SO}_2$  abundance at the cloud top as a function of the Venus diameter, the local time at the disk center and the longitude at the disk center (Fig. 14). There is no obvious relationship between any of these parameters with the  $\text{SO}_2$  abundance. It can be observed that, after 2023, the high values of  $\text{SO}_2$  are found when the nightside of Venus is observed (maximum diameter); however, the opposite is observed in January 2021, so no firm conclusion can be drawn.

Finally, we have used the long-term variations in the  $\text{SO}_2$  vmr at  $z = 57$  km, measured at  $18.9 \mu\text{m}$ , to study the evolution



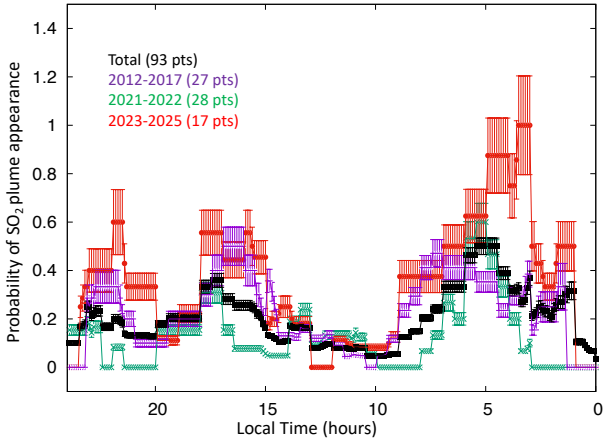
**Fig. 15.** Long-term variations in the  $\text{SO}_2(57 \text{ km})/\text{SO}_2(62 \text{ km})$  ratio between 2012 and 2025 (blue points), obtained from the disk-integrated  $\text{SO}_2$  mixing ratios derived at  $7.4 \mu\text{m}$  ( $z = 62$  km) and  $18.9 \mu\text{m}$  ( $z = 57$  km).

with time of the  $\text{SO}_2$  gradient within the cloud, between  $z = 57$  and  $z = 62$  km. Figure 15 shows the  $\text{SO}_2(57 \text{ km})/\text{SO}_2(62 \text{ km})$  ratio as a function of time. Not surprisingly, this ratio is usually higher than 1, which is expected as  $\text{SO}_2$  is known to decrease as the altitude increases. It can be seen that, in most of the cases, this ratio ranges between 0.9 and 2.2. The mean value of 1.6 corresponds to a scale height of about 10 km, i.e., larger than the scale height of about 5 km observed above the clouds. However, there are some noticeable exceptions. Values as high as 4 were observed in July–September 2018, and also in Autumn 2021, a period of very strong plume activity (E23). In contrast, values as low as 0.4–0.6 were observed occasionally between 2021 and 2024. These low values suggest an inversion of the  $\text{SO}_2$  vertical profile within the cloud; as was discussed above, this behavior is also suggested at the disk center on October 7, 2023.

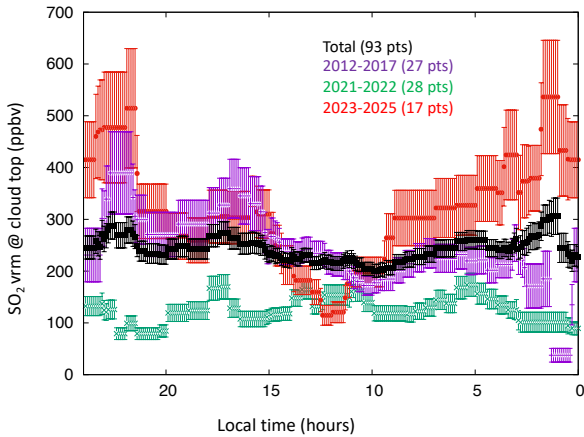
## 5. Variation of the $\text{SO}_2$ abundance as a function of local time

We have reconsidered our analysis of the  $\text{SO}_2$  variations at the cloud top as a function of the local time. As a reminder, our method consists in selecting one map per day and evaluating, on each map, the latitude, longitude and LT range of the  $\text{SO}_2$  plume (E23 and references therein). We then summed up all data from 2012 to 2022. Our previous study has shown that the  $\text{SO}_2$  plumes are mostly concentrated around the equator, while their distribution as a function of longitude has turned out to be inconclusive. For this reason, we concentrate here on the local time variations in  $\text{SO}_2$ . Previous studies have shown that the probability of appearance of the  $\text{SO}_2$  plumes was apparently higher around the terminators (E23).

Figure 16 shows the probability of  $\text{SO}_2$  plume appearance (independent of the  $\text{SO}_2$  abundance) as a function of local time, using the whole dataset (2012–2025, 93 pts), and Figure 17 shows the  $\text{SO}_2$  distribution as a function of local time, which can be compared to Venus Express and Akatsuki observations. In the present study, we have also taken into account the fact that the  $\text{SO}_2$  abundance as a function of time might be driven by different mechanisms and we have isolated three data sets corresponding to three different periods. The first period includes the 2012–2017 data (27 points), for which the  $\text{SO}_2$  vertical gradient was relatively uniform (Fig. 15). The second period includes the data taken in 2021–2022 (28 points), when the  $\text{SO}_2$  abundance was especially low (Fig. 12 and 13). The last period is the most



**Fig. 16.** Probability of  $\text{SO}_2$  appearance as a function of local time. Black: total (2012–2025, 93 points); violet: 2012–2017 (27 points); green: 2021–2022 (28 points); red: 2023–2025 (17 points). The error bar is proportional to  $n^{-0.5}$ , where  $n$  is the number of observations for which the local time is observed. The local hour scale increases from right to left, corresponding to the orientation on the Venus disk as seen by an external observer.



**Fig. 17.** Distribution of the  $\text{SO}_2$  vmr as a function of local time. Black: total (2012–2025, 93 points); violet: 2012–2017 (27 points); green: 2021–2022 (28 points); red: 2023–2025 (17 points). The error bar is proportional to  $n^{-0.5}$ , where  $n$  is the number of observations for which the local time is observed. The local hour scale increases from right to left, corresponding to the orientation on the Venus disk as seen by an external observer.

recent one (2023–2025, 17 points) which shows strong variations in the  $\text{SO}_2$  vmr.

It is interesting to note that the different data sets lead to significant differences. The probability of  $\text{SO}_2$  plume appearance shows, in all cases, a clear minimum around noon. Taking into account the whole data set, two maxima appear around 05:00 and 17:00, one hour before terminator. The  $\text{SO}_2$  vmr distribution shows even stronger differences between the different data sets. A maximum  $\text{SO}_2$  abundance appears in the afternoon (especially around 16:00) in 2012–2017, but there is no clear variation with local time in 2021–2022. In contrast, the last data set (2023–2025) exhibits a strong minimum around noon. These results are discussed in Section 6.

## 6. Discussion and conclusions

### 6.1. The $\text{SO}_2$ vertical distribution

When we started modeling the  $\text{SO}_2$  transitions at 7.4 and 18.9  $\mu\text{m}$ , we soon realized that these lines were broader than the  $\text{CO}_2$  and  $\text{H}_2\text{O}$  lines, which implied that a cutoff was needed in the  $\text{SO}_2$  vertical distribution a few kilometers above the clouds. In our latest calculations (E23), this cutoff, slightly depending on the thermal profile used, was fixed at  $z=70$  km. The relations between the  $\text{SO}_2/\text{CO}_2$  line depth ratios and the  $\text{SO}_2$  vmr were derived assuming this cutoff, and refer to the pressure levels of 100 mbar and 250 mbar, respectively.

The detection of the  $\text{SO}_2$  band at 8.6  $\mu\text{m}$  introduces a new constraint in the  $\text{SO}_2$  vertical distribution. Figures 3–5 and 9–11 show that we have difficulties to fit the three spectra simultaneously, both in the case of disk-integrated and located spectra. They also illustrate that our determination of the  $\text{SO}_2$  vmr at  $z=57$  km and  $z=62$  km is model dependent, as the amplitude of the  $\text{SO}_2$  absorption at 7.4 and 18.9  $\mu\text{m}$  depends upon the cutoff altitude assigned in our model. This means that the equations used in Section 3 to infer the  $\text{H}_2\text{O}$  and  $\text{SO}_2$  volume mixing ratios are adequate in the case of  $\text{H}_2\text{O}$  (because the  $\text{H}_2\text{O}$  vmr appears to be constant with height in the altitude range probed by TEXES), but is model-dependent in the case of  $\text{SO}_2$ , which exhibits significant variations as a function of altitude. Nevertheless, assuming for  $\text{SO}_2$  a constant altitude cutoff, as we did previously, allows us to determine what we call the  $\text{SO}_2$  vmr at 57 and 62 km, which we consider as representative indicators of the  $\text{SO}_2$  abundances at these altitudes. This simple method has allowed us to monitor the  $\text{SO}_2$  long-term evolution and to compare it to the temporal evolution of  $\text{H}_2\text{O}$ . It has also allowed us to build maps of the two species over the Venus disk, which have revealed the existence of the  $\text{SO}_2$  plumes and have been used to monitor their short-term evolution.

Ideally, an inversion algorithm should be used to retrieve the  $\text{H}_2\text{O}$  and  $\text{SO}_2$  vertical profiles. This analysis is beyond the scope of the present study, but will be considered in a future work. Here we limit ourselves to a few general comments. The strong modulation of the  $\text{SO}_2$  lines, in Figures 10 and 11, suggests that, at least in the two specific locations observed on October 7, 2025,  $\text{SO}_2$  is present above the cutoff level introduced at  $z=72$  km in our calculations. Another surprising feature is the very low  $\text{SO}_2$  vmr derived at  $z=57$  km at the disk center (Fig. 7). An inversion in the vertical profile of the  $\text{SO}_2$  vmr (being lower at 57 km than at 62 km), seems to be required to reconcile the data. It can be noted that this strange behavior is also observed in a few occasions on the disk-integrated  $\text{SO}_2$  vmr, especially in 2021–2023 (Fig. 15).

Another comment can be made about the parallel band structure observed on October 7, 2023. It can be seen from Figure 7 that the life time of this uncommon dynamical structure is exceptionally long, since it is continuously observed from October 3 to October 7, following the 4-day period rotation at the cloud top. In the past, we have seen only one example of a dynamical structure lasting more than a day, in November 7–9, 2021 (E23).

### 6.2. Long-term evolution of $\text{SO}_2$ and $\text{H}_2\text{O}$

Figures 12 and 13 illustrate the influence of the solar cycle on the long-term evolution of  $\text{SO}_2$  and  $\text{H}_2\text{O}$  at the cloud top of Venus. In particular, the increase in the solar irradiance in 1978–1981 is likely to be responsible for the sharp decrease of the  $\text{SO}_2$  vmr which does not necessarily require episodic volcanic activity, as

suggested previously by Esposito et al. (1984) and Marcq et al. (2013). The solar cycle also explains well the evolution of SO<sub>2</sub> and H<sub>2</sub>O between 2012 and 2023. However, the sharp decrease of SO<sub>2</sub> observed by Venus Express in 2006–2009 does not coincide with an increase in the solar activity, and the strong fluctuations of the SO<sub>2</sub> vmr since 2023 remain a mystery. These anomalies suggest that other mechanisms are involved, possibly associated with dynamical motions, sublimation/condensation processes or SO<sub>2</sub> gas-aerosol conversion.

### 6.3. SO<sub>2</sub> distribution as a function of local time

The changes observed in the SO<sub>2</sub> vmr distribution with local time between the three different period considered here (2012–2017, 2021–2022, 2023–2025, Fig. 17) could also be the signature of changes of regime. The SO<sub>2</sub> maximum observed in the afternoon in 2012–2017 is in good agreement with the UV measurements of Akatsuki which probe a slightly higher altitude, at about 70 km (Iwanaka et al. 2025) which were taken in 2016–2022. It is also in good agreement with the global circulation models (Stolzenbach et al. 2023). However, this trend is no more visible in the TEXES data after 2021. In 2021–2022, the SO<sub>2</sub> distribution as a function of local time is remarkably flat, while the 2023–2025 SO<sub>2</sub> distribution shows a strong minimum around noon, in better agreement with the SPICAV/Venus Express data taken earlier (2006–2015; Marcq et al. 2020). This temporal evolution of the SO<sub>2</sub> distribution as a function of local time is puzzling. It should be noted that the effects of local time and longitude are difficult to disentangle, and topographic effects (not considered in this analysis) could also possibly play a role in the SO<sub>2</sub> distribution along the equator.

### Data availability

The TEXES image cubes used in this paper are archived at the InfraRed Science Archive (IRSA), operated by the Infrared Processing and Analysis Center (IPAC) of the California Institute of Technology (<https://irsa.ipac.caltech.edu/frontpage/>) and can be downloaded at <https://irsa.ipac.caltech.edu/applications/irtf/>, specifying the Venus NAIF id 299 and the observing runs 2023A009 (March and July 2023), 2023B002 (October and December 2023), 2024A006 (February 2024), 2024B017 (January and March 2025). Data are available to the community after a 18-month proprietary period. The solar data are taken from the Solar Influences Data Analysis Center of the Royal Observatory of Belgium (<https://sidc.be/SILSO/datafiles#total>). The ephemeris are taken from the JPL Horizons system (<https://ssd.jpl.nasa.gov/horizons>).

**Acknowledgements.** TE, TKG and RG were visiting astronomers at the NASA Infrared Telescope Facility, which is operated by the University of Hawaii under Cooperative Agreement no. NNX-08AE38A with the National Aeronautics and Space Administration, Science Mission Directorate, Planetary Astronomy Program. We wish to thank the IRTF staff for the support of TEXES observations. This work was supported by the Programme National de Planétologie (PNP) of CNRS/INSU, co-funded by CNES. TKG acknowledges support of NASA Grant NNX14AG34G. TE and BB acknowledge support from CNRS. TW acknowledges support from the University of Versailles-Saint-Quentin and the European Commission Framework Program FP7 under Grant Agreement 606798 (Project EuroVenus). ML acknowledges funding from the European Union's Horizon Europe research and innovation program under the Marie Skłodowska-Curie grant agreement 101110489/MuSICA-V. WS acknowledges support from China Scholarship Council Fellowship.

### References

- Belyaev, D. A., Montmessin, F., Bertaux, J.-L., et al. 2012, *Icarus*, 217, 740  
 Bézard, B., & DeBergh, C. 2012, *J. Geophys. Res.*, 112, E04S07  
 Encrenaz, T., Greathouse, T. K., Richter, M. J., et al. 2013, *Astron. Astrophys.*, 559, A65  
 Encrenaz, T., Greathouse, T. K., Giles, R., et al. 2023, *Astron. Astrophys.*, 674, A199  
 Esposito, L. W. 1984, *Science*, 223, 1072  
 Esposito, L. W., Bertaux, J.-L., Krasnoposky, V., et al. 1997, Chemistry of lower atmosphere and clouds, in *Venus II: Geology, Geophysics, Atmosphere, and Solar Wind Environment*, eds. Bougher, S. W., Hunten, D. M., & Phillips, R. J., 415  
 Fedorova, A., Korablev, O., Vandaele, A.-C., et al. 2008, *J. Geophys. Res.*, 113, E00B25  
 Giles, R., Greathouse, T. K., Irwin, P., et al. 2022, *Icarus*, 387, 115187  
 Iwanaka, T., Imamura, T., Aoki, S., et al. 2025, *JGR Planets*, doi:10.1029/2024/E008775  
 Jacquinet-Husson, N., Scott, N., Chedin, A., et al. 2018, *J. Quant. Spectr. Rad. Transfer* 109, 1043  
 Krasnopolsky, V. A. 1986, *Photochemistry of the atmospheres of Mars and Venus* (New York: Springer Verlag)  
 Krasnopolsky, V. A. 2007, *Icarus*, 191, 25  
 Krasnopolsky, V. A. 2010, *Icarus*, 209, 314  
 Krasnopolsky, V. A., Belyaev, D. A., Gordon, I. A., et al. 2013, *Icarus*, 224, 57  
 Lacy, J. H., Richter, M. J., Greathouse, T. K., et al. 2002, *Pub. Astron. Soc. Pac.*, 114, 153  
 Marcq, E., Bertaux, J.-L., Montmessin, F., et al. 2013, *Nat. Geosci.*, 6, 25  
 Marcq, E., Mills, F. P., Parkinson, C. P., & Vandaele, A. C. 2018, *Space Sci. Rev.*, 214, 10  
 Marcq, E., Jessup, K. L., Baggio, L., et al. 2020, *Icarus*, 335, 11368  
 Mills, F. P., Esposito, L. W., & Yung, Y. K. 2007, in *Exploring Venus as a Terrestrial Planet*, Geophysical Monograph Series, 176, 73  
 Nakazawa, T., & Tanaka, M. 1982, *J. Quant. Spectr. Rad. Transfer*, 28, 409  
 Parkinson, C., Gao, P., Esposito, L., et al. 2015, *Plan. Space Sci.*, 112, 226  
 Rohlfs, K., & Wilson, T. L. 2004, *Tools for Radioastronomy*, 4th edn. (Berlin: Springer)  
 Stolzenbach, A., Lefèvre, F., Lebonnois, S., et al. 2023, *Icarus*, 395, 115447  
 Titov, D. V., Ignatiev, N. I., Mc Gouldrick, K., et al. 2018, *Space Sci. Rev.*, 214, 126  
 Vandaele, A.-C., Korablev, O., Belyaev, D., et al. 2017a, *Icarus*, 295, 16  
 Vandaele, A.-C., Korablev, O., Belyaev, D., et al. 2017b, *Icarus*, 295, 1  
 Zhang, K., Liang, M. C., & Mills, F. P. 2012, *Icarus*, 217, 714  
 Zasova, L. V., Moroz, V. I., Esposito, L. W., & Na, C. Y. 1993, *Icarus*, 105, 92